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# Hydrogeochemical and mineralogical characteristics related to heavy metal attenuation in a stream polluted by acid mine drainage: A case study in Dabaoshan Mine, China

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#### Abstract

Dabaoshan Mine, the largest mine in south China, has been developed since the 1970s. Acid mine drainage (AMD) discharged from the mine has caused severe environmental pollution and human health problems. In this article, chemical characteristics, mineralogy of ocher precipitations and heavy metal attenuation in the AMD are discussed based on physicochemical analysis, mineral analysis, sequential extraction experiments and hydrogeochemistry. The AMD chemical characteristics were determined from the initial water composition, water-rock interactions and dissolved sulfide minerals in the mine tailings. The waters, affected and unaffected by AMD, were Ca-SO<sub>4</sub> and Ca-HCO<sub>3</sub> types, respectively. The affected water had a low pH, high SO<sub>4</sub><sup>2-</sup> and high heavy metal content and oxidation as determined by the Fe<sup>2+</sup>/Fe<sup>3+</sup> couple. Heavy metal and SO<sub>4</sub><sup>2-</sup> contents of Hengshi River water decreased, while pH increased, downstream. Schwertmannite was the major mineral at the waste dump, while goethite and quartz were dominant at the tailings dam and streambed. Schwertmannite was transformed into goethite at the tailings dam and streambed. The sulfate ions of the secondary minerals changed from bidentate- to monodentate-complexes downstream. Fe-Mn oxide phases of Zn, Cd and Pb in sediments increased downstream. However, organic matter complexes of Cu in sediments increased further away from the tailings. Fe<sup>3+</sup> mineral precipitates and transformations controlled the AMD water chemistry.

2005).

Key words: acid mine drainage; Dabaoshan Mine; heavy metal attenuation; hydrogeochemical; PHREEQC

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### Introduction

The oxidation of sulfide minerals (e.g., pyrite) to produce acid mine drainage (AMD) at mine sites has been widely studied (Task, 1989) because AMD with high levels of sulfate and heavy metals at low pH (Mylona et al., 2000) decreases the biodiversity of water-ecosystems (Sydnor and Redente, 2002; Tordoff et al., 2000), contaminates the soil and plants, and endangers human health.

Dabaoshan Mine, the largest mine in south China, has been extensively mined since the 1970s, with massive tailings left behind and the surroundings have borne the effects of AMD contamination (Liu et al., 2009; Zhou et al., 2007). Production of AMD continues in the tailings reservoir (Bao et al., 2009; Lin et al., 2005b).

Contamination from Dabaoshan Mine has been of concern since cadmium pollution occurred in the Beijiang River (the main tributary of the Pearl River) in December 2005. The local government and the Dabaoshan Mining Corporation Limited have taken measures to reduce the impacts of the AMD. These measures include reinforcing plant contamination (Li et al., 2009; Lin et al., 2005a, 2005b; Liu et al., 2009; Zhou et al., 2007), toxicity to aquatic life (Chen et al., 2007; Lin et al., 2007), human health assessment of the residents (Bao et al., 2009; Zhuang et al., 2009a, 2009b) and the magnetic characteristics of contaminated soils (Zhou and Xia, 2010). However, they neglected to discuss the hydrogeochemical and secondary mineral characteristics of the waters affected by AMD, such as water chemical composition, secondary

the tailings reservoir, the revegetation of waste mining land and AMD treatment, but the problem still exists.

The literature reports that low pH, high-level heavy metal

content and intense rainfall were the main reasons for the

ineffectiveness of the measures (Sánchez-España et al.,

tal impact of AMD in this area, especially soil, water and

Most research groups have focused on the environmen-

mineral transformations and heavy metal natural attenuation in the AMD. These characteristics are important not only for establishing the solution chemistry of the AMD The environmental problems in the vicinity of the mine. but also for environmental remediation.

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are mainly caused by AMD pollution (Wu et al., 2009). The high levels of heavy metals are toxic to both aquatic life and terrestrial plants. Low pH values may damage the environmental enzymatic systems and decrease the respiratory action of plants. Removal of the heavy metals and raising the pH value of AMD-affected water are essential for environmental remediation. Secondary minerals (such as, Fe, Al and Mn oxyhydroxides) adsorb the heavy metals in AMD and the sorption process is essential for heavy metal attenuation in AMD (Bigham et al., 1996b; Fukushi et al., 2003; Kumpulainen et al., 2007; Munk et al., 2002; Ranville et al., 2004; Sánchez-España et al., 2005, 2006; Wu et al., 2009).

The sulfide tailings are classified into arkosite-type, carbonate-type, siliceous-type and silicate-type according to the gangue minerals. AMD produced by carbonate-type tailings are ignored based on the knowledge that limestone reacts with sulfide minerals and buffers the AMD. The processes are described by the following reactions (Holmström et al., 1999):

$$FeS_2 + 3.75O_2 + 3.5H_2O + 4CaCO_3 \longrightarrow Fe(OH)_3 + 4HCO_3^- + 2SO_4^{2+} + 4Ca^{2+}$$
(1)

Fe<sub>1-x</sub>S + (9-3x)/4O<sub>2</sub> + (5-3x)/2H<sub>2</sub>O + 2xCaCO<sub>3</sub> 
$$\longrightarrow$$
 (1-x)Fe(OH)<sub>3</sub> + 2(1-x)H<sup>+</sup> + SO<sub>4</sub><sup>2+</sup> + 2xCa<sup>2+</sup> + 2xHCO<sub>3</sub><sup>-</sup> (2)

where, x is in the range of 0–0.125.

However, AMD generated from carbonate-type tailings have been observed (Morin and Hutt, 1997). An armor

is formed as the carbonate minerals coat the secondary mineral precipitates, biofilms and bacteria, and the ratios of Reactions (1) and (2) decrease (Robbins et al., 1999). Therefore, the hydrogeochemical and secondary mineral characteristics of carbonate-type tailings are essential for understanding the AMD polluted areas.

The average contents of CaO and MgO of Dabaoshan Mine were 7.82 wt.% (n = 26) and 3.58 wt.% (n = 26) 26), respectively (Ge and Han, 1987). The tailings belong to the carbonate-type. However, a high level of heavy metal concentrations was determined in the Hengshi River, which is a tributary of the Beijiang River. Moreover, the Beijiang River is a primary water resource of Guangzhou. Therefore, to investigate the migration, transformation and attenuation of heavy metal in the Hengshi River is important both for water resource management and for AMD theory. To discuss the hydrogeochemical and secondary mineral characteristics and heavy metal attenuation of AMD in the vicinity of the Dabaoshan Mine, the following topics were investigated: (1) the physicochemical characteristics of the affected water; (2) the mineralogy of the ocher precipitates; and (3) natural heavy metal attenuation of AMD.

#### 1 Materials and methods

#### 1.1 Location and geological setting

Dabaoshan Mine (24°34′28″N, 113°43′42″E) (Fig. 1) is located in the north of Guangdong Province, south China. It has a subtropical humid monsoon climate. The average annual temperature is 20°C and rainfall is 1800 mm.

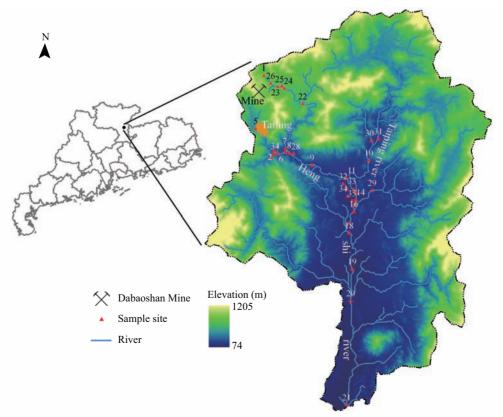


Fig. 1 Location and sample station map of the Dabaoshan Mine.

The Dabaoshan Mine belongs in the metallogenic province of polymetallic sulfide-pyrites in the north of Guangdong. The metallogenic province consists of western Guangxi, southern Hunan, and northern Guangdong. Several deep fractures pass through the area, and large-sized and medium-sized deposits are scattered in the fractures, and Dabaoshan Mine is one of the deposits. Iron, copper, zinc and lead are large-sized deposits in the mine and sulfur, tungsten and molybdenum are potentially available for industrial utilization.

The Dabaoshan Mountain is a huge intrusion of granitic rock. A volcanic eruption occurred in the Devonian period and dacite tuff and tuffaceous shale were widely spread in the ore belts. The deposit belongs to the class of volcanosedimentary formations and was formed about the middle-early Devonian period.

#### 1.2 Sample collection and preparation

Thirty five water samples in the vicinity of the Dabaoshan Mine were collected in June, 2009. The waters were classified into two groups based on both their hydrodynamic relationships and geological conditions. The first is affected water, which received the AMD directly, and the second is unaffected water, which was not influenced by AMD and is regarded as background water. All sample sites are shown in Fig. 1. Locations were established using a GARMIN global positioning system (GPS).

Nine sediments were collected during the water sampling period. Sediments included waste dump sediment (site 1), tailings dam sediment (site 5) and streambed sediment. Sediment sample locations were co-located with surface water sample sites. Streambed sediments were sampled along the Hengshi River. Site 18, the most distant collection point of sediment samples, was about 20 km away from the tailings dam (Fig. 1). At each sampling site, the top layer sediment (10 cm) was collected using a grab sampler, and a composite sample was formed by the mixing of 5 grab samples within 10 m<sup>2</sup> to represent each sampling site.

At each water sampling location, three sub-samples were collected across the river to form a composite sample to represent each sampling location. Water samples were filtered through 0.45  $\mu m$  cellulose filters in situ, and the following were added: (1) supra pure HCl for iron speciation analysis; (2) supra pure HNO<sub>3</sub> to keep pH < 2 for heavy metal and other trace element analyses; and (3) no additional treatment for analyses of major anions. Water samples were stored in high-density polyethylene bottles in an icebox. Sediment samples were stored in dark bottles in the icebox. All the samples were transported to the laboratory within 12 hr. Water samples were stored at 4°C before analysis, and sediment samples were freezedried immediately.

#### 1.3 Analysis procedures and methods

Temperature, pH, electrical conductivity (EC), oxidation-reduction potential (ORP), and total dissolved solids (TDS) were measured *in situ* with portable instruments (Orion 4-star, Thermo Scientific, USA).

Dissolved ion concentrations were measured using conventional methods. The major anions  ${\rm Cl}^-$ , and  ${\rm SO_4}^{2-}$  were analyzed with a Metrohm 732 ion chromatograph (Switzerland). Alkalinity was measured by an acid-base titration method with 0.025 mol/L HCl solution at pH of 4.4–4.5, within 24 hr after the samples were collected. The accuracy and precision were tested through triplicate analyses on selected samples.

Total Fe and  $Fe^{2+}$  were measured with 1, 10-Phenanthroline spectrophotometrically.  $Fe^{2+}$  was analyzed by colorimetry at 510 nm with a Shimadzu UV2450 spectrophotometer (Japan). The detection limit was 0.1 mg/L.

Heavy metals and other trace element concentrations were measured with inductively coupled plasma-optical emission spectrometry (ICP-OES) (Optima 5300DV, Perkin Elmer, USA) (Cánovas et al., 2007). K, Na, Ca and Mg were analyzed by Atomic Absorption Spectrometry (AAS) (Z-5000, Hitachi, Japan). Total As was determined by atomic fluorescence spectrometry (AFS-820, Beijing Titan, China). The accuracy and precision of the analytical methods were verified against certified reference materials: GSB04-1767-2004 (heavy metal and trace elements), GSBZ50020-90 (major cation) and GSBZ50004-88 (arsenic). The percent recoveries of elements were between 80% and 110%. A triplicate analysis was performed to evaluate the precision.

Sediment samples were pulverized before analysis. Mineralogy was characterized by powder X-ray diffraction (XRD), field emission scanning electron microscopy (FE-SEM) and Fourier transform infrared (FT-IR). The XRD analysis was carried out with an X-ray diffractometer (D/Max 2200 VPC, RIGAKU, Japan). The X-ray diffractometer was fitted with a 1.2 kW Cu Ka X-ray source. Diffractograms were collected in step-scan mode (0.02°) in 20 range of 3-80°. Patterns were interpreted with the aid of Scintag and MDI applications JADE search/match software and compared with reference patterns in the powder diffraction file (ICDD, 2002). The morphology of the secondary minerals was observed using FESEM (Quanta 400F, FEI/OXFORD/HKL, Dutch). The FT-IR spectra of the sediments were collected by the KBr pellet technique, and the spectra were collected in transmission mode in the 4000-400 cm<sup>-1</sup> range with a spectral resolution of 4 cm<sup>-1</sup>. The measurements were carried out on a FT-IR spectrometer (Nicolet iS10, Thermo Scientific, USA).

The classical 5-step Tessier sequential extraction procedure was used in this study (Table 1) (Tessier et al., 1979). The sequential extractions were performed in triplicate on 0.2 g of dried sediment. The reagents used during the extraction procedure were analysis grade and were prepared with Milli-Q water. Following each step, except for the residual, the extraction solutions were centrifuged at 5000 r/min for 30 min. The extraction solutions were filtered into sample vials and stored at 4°C before analysis. Heavy metals were measured by ICP-OES.

The PHREEQC geochemical modeling software package (v.2.15, February 5, 2008) (Parkhurst and Appelo, 1999) with wateq4f.dat was applied for the calculation

O.

Table 1	Sequential	extraction	procedure	used in	this study*
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Fraction	Extractant (dilution)	Procedure
Exchangeable fraction Adsorbed carbonates Fe-Mn oxides	1 mol/L MgCl <sub>2</sub> (pH = 7) 1 mol/L NaOAc (pH = 5) 0.04 mol/L NH <sub>2</sub> OH-HCl	Continuous shaking for 2 hr at room temperature Continuous shaking for 3 hr at room temperature Heat in water bath 96°C for 6 hr
Organic matter	0.04 mol/L HNO <sub>3</sub> -H <sub>2</sub> O <sub>2</sub> and 3.2 mol/L NH <sub>4</sub> OAc	30% H <sub>2</sub> O <sub>2</sub> added twice to samples with 0.04 mol/L HNO <sub>3</sub> shaking in water bath 85°C for 3 hr, cool to room temperature, then added 3.2 mol/L NH <sub>4</sub> OAc for 30 min
Residual	HCl-HNO <sub>3</sub> -HClO <sub>4</sub>	Sample digested with HCl, HNO <sub>3</sub> and HClO <sub>4</sub> in a microwave digester

<sup>\*</sup> Tessier et al., 1979.

of saturation indices (SI), activities and theoretical pE (Eh) values based on the Fe<sup>2+</sup>/Fe<sup>3+</sup> redox couple. The thermodynamic database of schwertmannite ( $\log k = 18.00 \pm 2.50$ ) was enlarged with the data from Bigham et al. (1996b).

#### 2 Results and discussions

#### 2.1 Chemical characteristics of water

The characteristics of surface water are shown in Table 2. High concentrations of TDS,  $SO_4^{2-}$ , Cu and Zn, with low Pb, Cd and As were the main characteristics of the affected water. The results were similar to prior studies (Allen et al., 1996; Kim and Chon, 2001; Sánchez-España et al., 2005; Wu et al., 2009).

The average concentration of  $HCO_3^-$  (44.06 mg/L) in affected water was significantly lower than that in unaffected water (71.75 mg/L) in the study area. The higher  $HCO_3^-$  samples were in the area adjacent to the karst strata, implying that carbonate rock dissolution was the main source of  $HCO_3^-$ .

The concentrations of major ions,  $HCO_3^-$ ,  $SO_4^{2-}$ ,  $Cl^-$ ,  $Ca^{2+}$ ,  $Mg^{2+}$ ,  $Na^+$  and  $K^+$  fluctuated in affected water. All median values were higher than average values. The data distribution of the samples showed that the concentrations of these ions changed systematically with distance from the AMD source.

Piper diagrams are widely used for classifying hydrogeochemical types of water by major ion analysis (Halima et al., 2009). The content of major ions plotted on a Piper diagram showed obvious variation, especially for anion content. The dominant cation, Ca<sup>2+</sup>, was abundant in both affected and unaffected water (Fig. 2).

The unaffected water samples were of the Ca-HCO $_3$  type. However, the samples collected from affected water were the Ca-SO $_4$  type. This reveals that the affected water received more SO $_4$ <sup>2-</sup> from the mine tailing.

# 2.2 pH, electrical conductivity (EC) and Eh

Generally, pH is the main factor influencing the chemical processes of AMD. The variations of chemical components in the affected and unaffected water were related to the pH. The highest pH value (7.09) was measured in unaffected water, and the lowest pH value (2.59) was detected in affected water (Table 2). In the study area, the pH values increased downstream of the Hengshi River, especially as an unaffected river flowed into the affected river. For

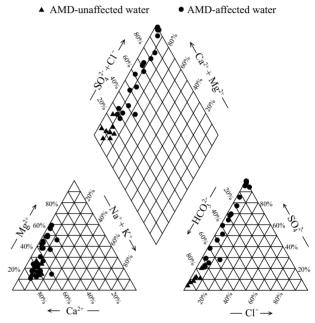


Fig. 2 Piper diagram showing the chemical composition of water samples in the study area.

example, pH increased substantially from 3.57 to 6.61 at Shangba village below the confluence of the unaffected river, the Taiping River, flowing into the Hengshi River. Lee and Kim (2008) observed a similar result.

The higher EC values in affected water revealed that more minerals were dissolved in affected water than in unaffected water. As shown in Fig. 3, the  ${\rm Ca^{2^+}}$ ,  ${\rm Mg^{2^+}}$  and  ${\rm SO_4^{2^-}}$  concentrations have good correlations with EC ( $R^2$  0.84, 0.94, and 0.93, respectively) indicating that EC might be used as an indicator of the contamination degree in AMD-affected water.

Water Eh is affected by several factors, such as dissolved oxygen, iron, chromium and manganese. However, the oxidation/reduction system of AMD is directly related to the oxidation of Fe<sup>2+</sup> (Sánchez-España et al., 2005). The calculated pE values based on the concentrations of Fe<sup>2+</sup> and Fe<sup>3+</sup> were well correlated ( $R^2 = 0.85$ , n = 17) with *in situ* measured Eh values (Fig. 4), which means the redox potential of the AMD was governed by Fe ions.

#### 2.3 Aluminum and iron

The average total Fe and  $Al^{3+}$  concentrations in affected water (12.95 mg/L and 13.07 mg/L, respectively) were much higher than those in unaffected water (0.16 mg/L and 0.1 mg/L, respectively). The Fe concentration was

Table 2 Statistical data for the geochemical parameters of the waters of Dabaoshan Mine in Guangdong, China

	Affected water $(n = 27)$			Unaffected water $(n = 8)$				
	Max	Min	Mean	Median	Max	Min	Mean	Median
pH	7.03	2.59	5.13	5.52	7.09	6.06	6.55	6.52
Eh (mV)	691.80	283.40	480.82	391.85	520.30	275.30	352.29	323.40
EC (μS/cm)	3090.00	25.35	698.28	342.50	596.00	40.30	194.50	111.40
TDS (mg/L)	1513.00	12.00	342.15	167.50	292.00	20.00	95.33	55.00
Cl <sup>-</sup> (mg/L)	13.96	0.19	1.90	0.85	2.15	0.19	0.82	0.47
$SO_4^{2-}$ (mg/L)	2108.00	2.89	295.37	77.97	19.28	2.18	6.58	4.61
$HCO_3^-$ (mg/L)	130.61	3.06	44.06	26.62	308.04	21.26	71.75	45.10
K <sup>+</sup> (mg/L)	24.27	0.17	3.58	1.65	4.50	0.24	1.33	1.22
Na <sup>+</sup> (mg/L)	9.27	0.13	2.30	1.32	7.29	0.34	1.83	1.33
Ca <sup>2+</sup> (mg/L)	490.00	0.66	67.59	35.61	78.42	5.13	19.26	8.97
$Mg^{2+}$ (mg/L)	201.60	0.46	35.81	8.74	12.00	1.07	3.17	1.88
$Cu^{2+}$ (mg/L)	7.67	0	1.74	0.29	0.01	0	0	0
$Pb^{2+}$ (mg/L)	0.70	0.02	0.28	0.19	bd	bd	bd	bd
Cd <sup>2+</sup> (mg/L)	0.25	0	0.09	0.07	bd	bd	bd	bd
$Zn^{2+}$ (mg/L)	34.64	0	6.85	1.30	0.76	0.02	0.16	0.05
Total Fe (mg/L)	115.64	0.02	12.95	2.54	0.40	0.01	0.16	0.09
Fe <sup>2+</sup> (mg/L)	68.62	0.01	6.21	1.17	0.14	0.01	0.07	0.07
$Al^{3+}$ (mg/L)	62.61	0	13.07	2.51	0.19	0.03	0.10	0.09
Total As (mg/L)	0.03	0	0.01	0.01	0.02	0	0.01	0
Mn <sup>2+</sup> (mg/L)	29.32	0	7.01	2.23	0.03	0	0.02	0.02
Sr <sup>2+</sup> (mg/L)	0.43	0	0.08	0.06	0.16	0.01	0.04	0.02

TDS: total dissolved solid; bd: below detection limits.

Affected water sample were collected from sites 1, 2, 4–9, 11–23, 25, 26, 28, 32, 33 and 35; unaffected water sample were collected from sites 3, 10, 24, 27, 29-31 and 34.

positively related with the Al concentration ( $R^2 = 0.86$ , n =

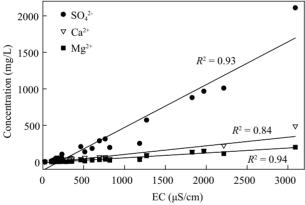
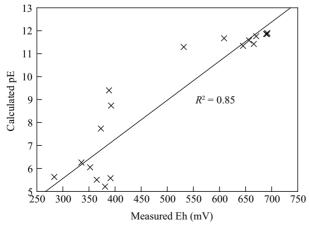


Fig. 3 Correlation plots of electrical conductivity (EC) vs.  $Ca^{2+}$ ,  $Mg^{2+}$  and  $SO_4^{2-}$  concentration in the affected water of Dabaoshan Mine.

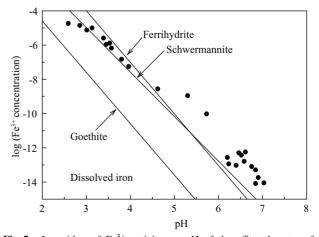


**Fig. 4** Correlations between measured Eh and calculated pE. The measured Eh values were obtained on-site, whereas the calculated pE were computed with PHREEQC v. 2.15, by introducing measured  $Fe^{2+}$  and  $Fe^{3+}$  concentrations.

27) in affected water, suggesting that Fe and Al in affected water originated from the minerals of the Dabaoshan Mine.

Fe exists as different species at different pH. Figure 5 shows that the concentration of  $Fe^{3+}$  in affected water decreases with rising pH, and the solubility of goethite and ferrihydrite have similar tendencies. The  $Fe^{3+}$  concentrations of all samples were higher than the solubility line of goethite at all pH, but below the solubility line of ferrihydrite at pH < 4.5, and above the line at pH > 4.5. This means that the deposition of Fe in affected water was controlled by pH.

Schwertmannite (ideal formula:  $Fe_8O_8(OH)_6SO_4$ ) proved to be very important for iron precipitation in AMD (Bigham et al., 1996a; Eskandarpour et al., 2008; Sánchez-España et al., 2005, 2006, 2007, 2008; Yu et al., 1999). The calculation of saturation indices revealed that sediment suspensions were supersaturated with



**Fig. 5** Logarithm of  $Fe^{3+}$  activity vs. pH of the affected water of the Dabaoshan Mine with goethite, ferrihydrite and schwertmannite solubility lines (Bigham et al., 1996b).

respect to goethite and schwertmannite; the waters are undersaturated below pH 3.2, at saturation between pH 3.2 and 4, and supersaturated above pH 4 (Fig. 5). Thus, the precipitate of schwertmannite may govern the activities of Fe<sup>3+</sup> ions above pH 4 and Fe ions could be rapidly removed from the AMD as Fe<sup>3+</sup> oxyhydroxide precipitates with increasing pH.

#### 2.4 Trace elements

Sediment is an important factor as AMD is discharged at mining sites. Sediment geology and mineralogy are helpful in understanding the environmental activities of heavy metals and trace elements in AMD.

Cadmium exists in sphalerite accompanied by Zn, and the geochemical characteristics of Cd are similar to Zn. The concentrations of Zn and Cd in affected water were well correlated ( $R^2=0.97,\,n=27$ ). Ullrich et al. (1999) and Shikazono et al. (2008) reported that Cd and Zn were always found together in nature and their concentrations were closely correlated. The correlations between Al and other metal ions are also significant, e.g. Al:Zn ( $R^2=0.98$ ), Al:Cu ( $R^2=0.94$ ), Al:Pb ( $R^2=0.97$ ), Al:Cd ( $R^2=0.97$ ) and Al:Mn ( $R^2=0.85$ ). It is possible that these elements originated from the same minerals at the mine site.

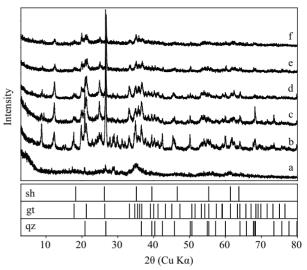
Another significant geochemical characteristic is the association of potassium and strontium ( $R^2 = 0.90$ , n = 27) in affected water. It may be inferred from this correlation that strontium and potassium originated from the same minerals.

#### 2.5 Mineralogy of AMD sediments

The most widespread feature of the study area is the ocher sediments on riverbanks affected by AMD. These sediments consist of Fe-phases precipitated from the Fe dissolved in AMD and coming from pyrite at the mine site. The minerals of AMD sediment have been well characterized in many mine areas all over the world by different techniques (SEM, XRD and FT-IR). The minerals consist of tiny particle size (from  $\mu$ m to nm in diameter), amorphous oxyhydroxides and oxyhydroxysulfates with fibrous to spherical appearance, such as schwertmannite and goethite.

Ocher sediments of Fe-minerals were found at the waste dumps (site 1), tailings dam (site 5) and river (sites 8, 9, 16 and 18) in the Dabaoshan Mine area. The secondary minerals are formed as the AMD leaches from the mine tips. Soil-free or loosely held modes can be observed in the entire system.

Schwertmannite is the dominant mineral at site 1 (waste dumps) as shown by XRD analysis (Fig. 6 line a), which due to the continuous injection and slow flow of AMD. Goethite and quartz with trace amounts of schwertmannite are present at sites 5 (tailings dam), 8, 9, 16 and site 18 (Fig. 6 line b) (Hengshi River). The presence of goethite together with trace schwertmannite indicates that schwertmannite is metastable with respect to goethite in these systems. The transformation of schwertmannite to goethite has been observed both under laboratory conditions (Acero et al., 2006; Bigham et al., 1996b; Jönsson et al., 2005;



**Fig. 6** X-ray powder diffraction pattern of precipitates from the waste dump (site 1 line a) and tailings dam (site 5, line b) and Hengshi creek; line c: site 8; line d: site 9; line e: site 16; line f: site 18. gt: goethite, sh: schwertmannite, qz: quartz.

Kawano and Tomita, 2001) and under natural conditions (Asta et al., 2010; Gagliano et al., 2004; Kumpulainen et al., 2007; Pérez-López et al., 2011). The minerals changed in the tailings dam and riverbanks. This might be due to the warm weather in the study area and the fact that schwertmannite had enough time to be transformed into goethite. The transformation occurs via the overall reaction (Bigham et al., 1996b):

$$Fe_8O_8(OH)_{5.5}(SO_4)_{1.25(s)} + 2.5H_2O_{(l)} = 8FeOOH_{(s)} + 2.5H_{(aq)}^+ + 1.25SO_4^{2-}_{(aq)}$$
(3)

SEM analysis of site 1 showed that schwertmannite formed spherical particles of  $1{\text -}2~\mu\text{m}$  diameter (Fig. 7) that were associated in larger aggregates. Both spherical and fibrous particles observed at site 5 indicated that schwertmannite and goethite coexisted. Characteristic filamentous features were observed for the schwertmannite spherical particles in site 1. Spherical phases were present at sites 8 and 9. The SEM revealed a low crystalline phase at sites 16 and 18 and ultrafine particles adhered to the grains.

FT-IR has been widely used to analyze the minerals generated from AMD. The studies have mainly focused on the minerals generated under conditions of different pH (Boily et al., 2010; Jönsson et al., 2005; Peak et al., 1999; Zhang and Peak, 2007), temperature (Boily et al., 2010) and depth (Gagliano et al., 2004). However, few have studied the FT-IR spectra of minerals generated at varying distance from the mine tailing under natural conditions.

The FT-IR analysis of site 1 exhibited the characteristic OH deformations (700 and 850 cm $^{-1}$ ) (Fig. 8a) and stretches (3400 cm $^{-1}$ ) (not shown) of schwertmannite, which were consistent with the XRD and SEM analysis. The  $\nu_{so}$  modes of schwertmannite consisted of a broad triply degenerate  $\nu_3$  band at 1128 cm $^{-1}$  with shoulders at 1040 and 1205 cm $^{-1}$ , a  $\nu_1$  fundamental band of the  $SO_4{}^{2-}$  stretch at 980 cm $^{-1}$ , and a  $\nu_4$  bending band at 610 cm $^{-1}$ . The spectra implied that bidentate sulfate complexes of  $C_{2\nu}$  symmetry occur on the schwertmannite surface



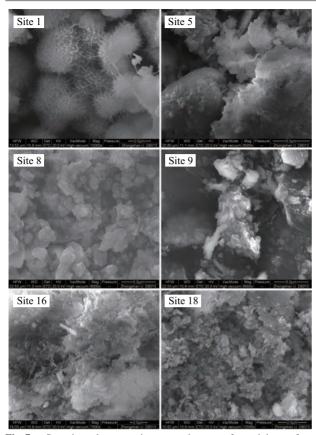
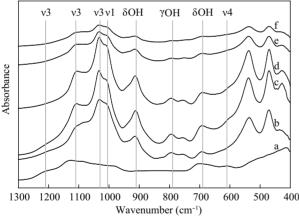


Fig. 7 Scanning electron microscopy images of precipitates from different sites.



**Fig. 8** FT-IR spectra of precipitates from the waste dump (site 1, line a) and tailings dam (site 5, line b) and Hengshi creek: (line c) site 8; (line d) site 9; (line e) site 16; and (line f) site 18.

(Nakamoto, 2009; Peretyazhko et al., 2009). The spectra of the sediment samples collected from the tailings dam (site 5) and Hengshi River (sites 8, 9, 16 and 18) were different from that of the waste dumps (site 1). The  $v_1$  band becomes active at  $1005 \text{ cm}^{-1}$ , while the  $v_3$  band split into 1106, 1032, and  $1208 \text{ cm}^{-1}$  at sites 5 and 8 (Fig. 8 line b, c). The  $1208 \text{ cm}^{-1}$  band vanished at sites 9, 16, and 18 (Fig. 8 line d–f), indicating that the bidentate  $C_{2\nu}$  complexes were transformed to monodentate complexes of  $C_{3\nu}$  symmetry on the mineral surface downstream in the river. Peretyazhko et al. (2009) reported a conversion of bidentate- to monodentate-bound sulfate complexes accompanied by sulfate desorption in AMD minerals. The

formation of goethite and other minerals, as influenced by the presence of co-precipitated silicate, showed broad FT-IR bands at 538 and 469 cm<sup>-1</sup> (silicate bending region) (Nakamoto, 2009; Vempati and Loeppert, 1989). The quartz present in the streambed sediments was confirmed by XRD analysis, as shown in Fig. 6.

The results indicated that fresh sediments at site 1 (waste dumps) were composed of metastable schwertmannite. The schwertmannite in sediments of the Hengshi River gradually transformed at distances over short time scales into goethite, and the quantities of goethite increased with the distance away from the tailing.

#### 2.6 Sequential extraction of sediment

The newer revised sequential extraction methods (Caraballo et al., 2009; Dold, 2003) used in sulfide wastes and in acid mine water passive remediation systems were developed after Tessier et al. (1979) introduced the five step sequential extraction method. To compare our results to other studies, we selected the classical Tessier 5-step sequential extraction method in this article. Five phases were extracted from the sediment samples (Fig. 9): exchangeable, carbonates, Fe-Mn oxides, organic matter and residual. The exchangeable phase contained the heavy metals adsorbed on the major constituents of the sediment. The carbonates fraction showed that some heavy metals were associated with sediment carbonate. This fraction was liberated with a pH change. The Fe-Mn oxides phase was cemented or coated on the particles. These phases were unstable under anoxic conditions. The organic matter phase consists of heavy metals bound to various organic matter forms. The residual fraction consists of the heavy metals held within the mineral crystal structures (Ranville et al., 2004; Tessier et al., 1979). Carbonates and Fe-Mn oxide extractions were considered chemically mobile. These phases could liberate heavy metals to the environment as the geochemical conditions change (Ranville et al., 2004).

Heavy metals show different affinities within the sequential extraction phases. Pb, Zn and Cu have a high affinity for absorption on Fe-hydroxides (Lee et al., 2002; Lottermoser et al., 1999; Ranville et al., 2004). Cu in the bioavailable phase (organic matter) increased after the AMD confluence with other creeks (Ranville et al., 2004), and the mineral form was Cu<sub>2</sub>(OH<sub>2</sub>)CO<sub>3</sub> as predicted by Chapman et al. (1983). Cd was primarily found in the silicate (residual) and Fe oxide phases (Ranville et al., 2004). The sequential extraction results are shown in Fig. 9. The exchangeable fraction of Cu, Zn, Pb and Cd decreased downstream in the river. The carbonate phases of Cu, Zn, Pb and Cd increased further away from the mine site. The Fe-Mn oxide phases of Zn, Cd and Pb increased, while Cu decreased except at site 9, downstream in the river. Organic matter phases of Zn and Cd decreased, but there was an increase in Cu downstream. The organic matter phase of Pb increased before site 9 and then decreased. The residual phase of the four elements dominated in the Hengshi river sediments. Except for the residual phase, the other four phases were different for Cu, Zn, Pb and Cd. The Fe-Mn

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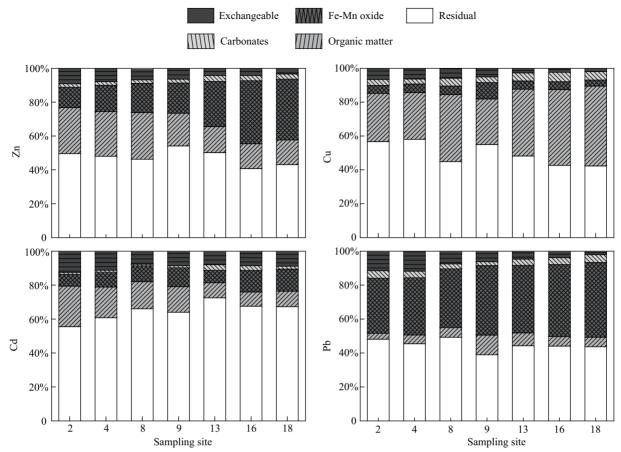


Fig. 9 Sequential extractions of the sediments of Hengshi River.

oxide phase was dominant for Pb, and carbonate was the major phase for Cu. The carbonate phase was dominant at site 2 (upstream) while the Fe-Mn was the main oxide phase at site 18 (downstream) for Zn.

# 2.7 Attenuation of heavy metals in AMD

AMD systems are complex and a combination of processes exists. Heavy metals in AMD can be removed by adsorption, deposition, co-precipitation and bioremediation (Neculita et al., 2007). The adsorption, deposition and co-precipitation are physicochemical processes in the systems, and the pH value affects their removal efficiency of heavy metals. As the unpolluted river inflow enters the AMD-affected river, the pH, ORP, TDS and other indexes of the systems would be changed. Moreover, the unpolluted river would dilute the systems and accelerate the reactions in the systems.

Attenuation of heavy metals in AMD has been proven by the variation in content of heavy metals and  $SO_4^{2-}$ . Heavy metal and  $SO_4^{2-}$  concentrations in the Hengshi River were above background values when the AMD flowed into the river, and decreased after the Taiping River inflow. At the confluence, the pH increased (from 3.57 to 6.61), and EC (from 693 to 105  $\mu$ S/cm), ORP (from 665.8 to 283.4 mV), HCO<sub>3</sub><sup>-</sup> (from 0 to 20.91 mg/L),  $SO_4^{2-}$  (from 287.25 to 10.83 mg/L) and heavy metals (from 1.58 to 0.01 mg/L for Cu, from 0.13 mg/L to below detection limits for Pb, from 0.04 mg/L to below detection limits for Cd, from 8.05 to 0.08 mg/L for Zn and from 0.01 mg/L to

below detection limits for As) decreased. The variation in concentrations of heavy metals and other ions in the Hengshi River may be explained by the secondary Fe or Al-oxyhydroxide minerals' selective adsorption. Removal of heavy metals by adsorption onto these minerals has been established under both field and laboratory conditions. Several studies revealed that Cu, Zn and Pb are adsorbed on the Fe hydroxide surface (Chapman et al., 1983; Dzombak and Morel, 1990; Johnson, 1986). Al hydroxides or hydroxysulfates can adsorb Pb, Cu, Zn and Ni (Munk et al., 2002). The maximum sorption of As occurred at pH 3-7, and declined with the increase of pH (Adriano, 2001). Fukushi et al. (2003) reported that As was rapidly scavenged from drainage water in abandoned arsenic mine dumps by sorption on hydrous iron oxides. Fe oxyhydroxide precipitation played an important role in the removal of heavy metals by adsorption and coprecipitation (Benjamin, 1983; Stumm and Sulzberger, 1992).

AMD could be neutralized by the dissolution of carbonate minerals, and the mine tailing belongs to the carbonate-type in the study area. The AMD of the Dabaoshan mine is initially strong acidic because H<sup>+</sup> ions liberated by sulfide oxidation are not consumed by the dissolution of carbonate minerals immediately. Limestone is widespread in the study area. When the AMD moves away from the mine tailing, the dissolution of carbonate minerals takes place, and the acidity produced by the sulfide minerals is buffered by the carbonate minerals. This

results in increase of the pH and decrease of heavy metal concentrations in water, and transformation of secondary minerals in sediment.

Due to the natural attenuation processes, such as secondary minerals adsorption and co-precipitation, carbonate minerals buffering, and unaffected water dilution, the hydrogeochemical characteristics of the affected water and the unaffected water become similar downstream in the Hengshi River.

#### 3 Conclusions

The chemical characteristics of affected water in the vicinity of the Dabaoshan Mine are controlled by sulfide minerals in spite of the AMD belonging to the carbonate-type. AMD is characterized by low pH (5.13) and high concentrations of heavy metals (1.74 mg/L Cu, 0.28 mg/L Pb, 0.09 mg/L Cd and 6.85 mg/L Zn) and SO<sub>4</sub><sup>2-</sup> (295.37 mg/L). In affected water, Ca<sup>2+</sup> and SO<sub>4</sub><sup>2-</sup> were the major ions, whereas Ca<sup>2+</sup> and HCO<sub>3</sub><sup>-</sup> dominated in unaffected water.

The good correlation ( $R^2 = 0.85$ , n = 17) between calculated pE and in situ measured Eh, suggests that the Fe<sup>2+</sup>/Fe<sup>3+</sup> couple determined the redox potential of the affected water. The pH increased, while the heavy metals and SO<sub>4</sub><sup>2-</sup> concentrations decreased downstream, especially at the unaffected river confluence. SEM, XRD, FT-IR, sequential extraction experiment and PHREEQC simulation results showed that secondary Fe minerals precipitated in the affected water, which was confirmed by the observation of ocher precipitations in the field. Schwertmannite was the major mineral at the waste dump, while goethite was dominant at the tailings dam and riverbank. The bidentate complexes of C<sub>2v</sub> symmetry transformed to monodentate complexes of C<sub>3v</sub> symmetry in the sulfate radical of secondary minerals, downstream. Fe-Mn oxide phases of Zn, Cd and Pb in sediments increased downstream, however, the organic matter complexes of Cu in sediments increased further away from the AMD sources.

Heavy metal attenuations are very complex processes in AMD. The heavy metal adsorption or co-precipitation on other minerals (such as: Al or Mn minerals) requires further study in this region. Moreover, a detailed research study of the mineral transformations should also be initiated.

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