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Highlight articles

- 129 Rice: Reducing arsenic content by controlling water irrigation
Ashley M. Newbigging, Rebecca E. Paliwoda and X. Chris Le
- 132 Apportioning aldehydes: Quantifying industrial sources of carbonyls
Sarah A. Styler

Review articles

- 30 Application of constructed wetlands for wastewater treatment in tropical and subtropical regions (2000-2013)
Dong-Qing Zhang, K.B.S.N. Jinadasa, Richard M. Gersberg, Yu Liu, Soon Keat Tan and Wun Jern Ng
- 47 Stepwise multiple regression method of greenhouse gas emission modeling in the energy sector in Poland
Alicja Kolasa-Wiecek
- 113 Mini-review on river eutrophication and bottom improvement techniques, with special emphasis on the Nakdong River
Andinet Tekile, Ilho Kim and Jisung Kim

Regular articles

- 1 Effects of temperature and composite alumina on pyrolysis of sewage sludge
Yu Sun, Baosheng Jin, Wei Wu, Wu Zuo, Ya Zhang, Yong Zhang and Yaji Huang
- 9 Numerical study of the effects of local atmospheric circulations on a pollution event over Beijing-Tianjin-Hebei, China
Yucong Miao, Shuhua Liu, Yijia Zheng, Shu Wang and Bicheng Chen, Hui Zheng and Jingchuan Zhao
- 21 Removal kinetics of phosphorus from synthetic wastewater using basic oxygen furnace slag
Chong Han, Zhen Wang, He Yang and Xiangxin Xue
- 55 Abatement of SO₂-NO_x binary gas mixtures using a ferruginous active absorbent: Part I. Synergistic effects and mechanism
Yinghui Han, Xiaolei Li, Maohong Fan, Armistead G. Russell, Yi Zhao, Chunmei Cao, Ning Zhang and Genshan Jiang
- 65 Adsorption of benzene, cyclohexane and hexane on ordered mesoporous carbon
Gang Wang, Baojuan Dou, Zhongshen Zhang, Junhui Wang, Haier Liu and Zhengping Hao
- 74 Flux characteristics of total dissolved iron and its species during extreme rainfall event in the midstream of the Heilongjiang River
Jiunian Guan, Baixing Yan, Hui Zhu, Lixia Wang, Duian Lu and Long Cheng
- 81 Sodium fluoride induces apoptosis through reactive oxygen species-mediated endoplasmic reticulum stress pathway in Sertoli cells
Yang Yang, Xinwei Lin, Hui Huang, Demin Feng, Yue Ba, Xuemin Cheng and Liuxin Cui
- 90 Roles of SO₂ oxidation in new particle formation events
He Meng, Yujiao Zhu, Greg J. Evans, Cheol-Heon Jeong and Xiaohong Yao
- 102 Biological treatment of fish processing wastewater: A case study from Sfax City (Southeastern Tunisia)
Meryem Jemli, Fatma Karray, Firas Feki, Slim Loukil, Najla Mhiri, Fathi Aloui and Sami Sayadi

CONTENTS

- 122 Bioreduction of vanadium (V) in groundwater by autohydrogentrophic bacteria: Mechanisms and microorganisms
Xiaoyin Xu, Siqing Xia, Lijie Zhou, Zhiqiang Zhang and Bruce E. Rittmann
- 135 Laccase-catalyzed bisphenol A oxidation in the presence of 10-propyl sulfonic acid phenoxazine
Rūta Ivanec-Goranina, Juozas Kulys, Irina Bachmatova, Liucija Marcinkevičienė and Rolandas Meškys
- 140 Spatial heterogeneity of lake eutrophication caused by physiogeographic conditions: An analysis of 143 lakes in China
Jingtao Ding, Jinling Cao, Qigong Xu, Beidou Xi, Jing Su, Rutai Gao, Shouliang Huo and Hongliang Liu
- 148 Anaerobic biodegradation of PAHs in mangrove sediment with amendment of NaHCO_3
Chun-Hua Li, Yuk-Shan Wong, Hong-Yuan Wang and Nora Fung-Yee Tam
- 157 Achieving nitrification at low temperatures using free ammonia inhibition on *Nitrobacter* and real-time control in an SBR treating landfill leachate
Hongwei Sun, Yongzhen Peng, Shuying Wang and Juan Ma
- 164 Kinetics of Solvent Blue and Reactive Yellow removal using microwave radiation in combination with nanoscale zero-valent iron
Yanpeng Mao, Zhenqian Xi, Wenlong Wang, Chunyuan Ma and Qinyan Yue
- 173 Environmental impacts of a large-scale incinerator with mixed MSW of high water content from a LCA perspective
Ziyang Lou, Bernd Bilitewski, Nanwen Zhu, Xiaoli Chai, Bing Li and Youcai Zhao
- 180 Quantitative structure-biodegradability relationships for biokinetic parameter of polycyclic aromatic hydrocarbons
Peng Xu, Wencheng Ma, Hongjun Han, Shengyong Jia and Baolin Hou
- 191 Chemical composition and physical properties of filter fly ashes from eight grate-fired biomass combustion plants
Christof Lanzerstorfer
- 198 Assessment of the sources and transformations of nitrogen in a plain river network region using a stable isotope approach
Jingtao Ding, Beidou Xi, Qigong Xu, Jing Su, Shouliang Huo, Hongliang Liu, Yijun Yu and Yanbo Zhang
- 207 The performance of a combined nitrification-anammox reactor treating anaerobic digestion supernatant under various C/N ratios
Jian Zhao, Jiane Zuo, Jia Lin and Peng Li
- 215 Coagulation behavior and floc properties of compound bioflocculant-polyaluminum chloride dual-coagulants and polymeric aluminum in low temperature surface water treatment
Xin Huang, Shenglei Sun, Baoyu Gao, Qinyan Yue, Yan Wang and Qian Li
- 223 Accumulation and elimination of iron oxide nanomaterials in zebrafish (*Danio rerio*) upon chronic aqueous exposure
Yang Zhang, Lin Zhu, Ya Zhou and Jimiao Chen
- 231 Impact of industrial effluent on growth and yield of rice (*Oryza sativa* L.) in silty clay loam soil
Mohammad Anwar Hossain, Golum Kibria Muhammad Mustafizur Rahman, Mohammad Mizanur Rahman, Abul Hossain Molla, Mohammad Mostafizur Rahman and Mohammad Khabir Uddin
- 241 Molecular characterization of microbial communities in bioaerosols of a coal mine by 454 pyrosequencing and real-time PCR
Min Wei, Zhisheng Yu and Hongxun Zhang
- 252 Risk assessment of *Giardia* from a full scale MBR sewage treatment plant caused by membrane integrity failure
Yu Zhang, Zhimin Chen, Wei An, Shumin Xiao, Hongying Yuan, Dongqing Zhang and Min Yang
- 186 Serious BTEX pollution in rural area of the North China Plain during winter season
Kankan Liu, Chenglong Zhang, Ye Cheng, Chengtang Liu, Hongxing Zhang, Gen Zhang, Xu Sun and Yujing Mu

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Assessment of the sources and transformations of nitrogen in a plain river network region using a stable isotope approach

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ABSTRACT

The great spatial and temporal variability in hydrological conditions and nitrogen (N) processing introduces large uncertainties to the identification of N sources and quantifying N cycles in plain river network regions. By combining isotopic data with chemical and hydrologic measurements, we determined the relative importance of N sources and biogeochemical N processes in the Taige River in the East Plain Region of China. The river was polluted more seriously by anthropogenic inputs in winter than in summer. Manure and urban sewage effluent were the main nitrate (NO_3^-) sources, with the nitrification of N-containing organic materials serving as another important source of NO_3^- . In the downstream, with minor variations in hydrological conditions, nitrification played a more important role than assimilation for the decreasing ammonium ($\text{NH}_4^+\text{-N}$) concentrations. The N isotopic enrichment factors (ϵ) during NH_4^+ utilization ranged from -13.88% in March to -29.00% in July. The ratio of the increase in $\delta^{18}\text{O}$ and $\delta^{15}\text{N}$ of river NO_3^- in the downstream was 1.04 in January and 0.92 in March. This ratio indicated that NO_3^- assimilation by phytoplankton was responsible for the increasing $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values of NO_3^- in winter. The relationships between $\delta^{15}\text{N}$ of particulate organic nitrogen and isotopic compositions of dissolved inorganic nitrogen indicated that the phytoplankton in the Taige River probably utilized NH_4^+ preferentially and mainly in summer, while in winter, NO_3^- assimilation by phytoplankton was dominant.

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Introduction

Nitrogen (N) contamination in freshwater systems is a worldwide environmental problem. High levels of N not only pose a potential threat to human health but also stimulate water eutrophication (Kendall, 1998; Lee et al., 2008). External supplies of N to river water are generally associated with inputs from anthropogenic activities (e.g., municipal and industrial wastewater, livestock wastewater, atmospheric

deposition and agricultural runoff). N usually undergoes numerous transformations, such as nitrification, assimilation, and denitrification; the occurrence and extent of these processes depend on a variety of environmental variables (Middelburg and Nieuwenhuize, 2001). Identification of the sources and forms of N, as well as their changes with time, is an important step in improving the management practices associated with maintaining water quality in rivers (Sugimoto et al., 2011; Li et al., 2010).

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Many studies have shown that stable isotope techniques are useful in understanding N sources and cycling in terrestrial and aquatic ecosystems (Wells and Krothe, 1989; Feast et al., 1998; Mayer et al., 2002; Curt et al., 2004). Dual isotope analysis of nitrate ($\delta^{15}\text{N}-\text{NO}_3^-$ and $\delta^{18}\text{O}-\text{NO}_3^-$) has been frequently used to differentiate NO_3^- sources because of the distinct isotopic characteristics of the main NO_3^- sources, such as rain, chemical fertilizers, and NO_3^- derived from nitrification (Xue et al., 2009; Nestler et al., 2011). To provide valuable information for identifying NO_3^- sources and the processes during NO_3^- retention and transport in watersheds, scientists have used this dual-isotope technique in combination with physicochemical water quality data, precipitation data, hydrological data, land use data, and other tracers (e.g., boron isotopes, chloride) in forested (Burns et al., 2009; Piatek et al., 2009), agricultural (Deutsch et al., 2006; Johannsen et al., 2008), urban, and mixed watersheds (Kaushal et al., 2011; Chen et al., 2012).

Zhushan Bay is a semi-closed lacustrine bay in the north of Taihu Lake, the third largest freshwater lake in the East Plain Region in China (Fig. 1). Zhushan Bay suffers from steadily worsening eutrophication caused by the inflow of nutrients from the Taige River (Yan et al., 2011; Zhang et al., 2011). The river catchment basin is flat, with intersecting rivers and water flowing slowly (annual mean flow velocity is 0.13 m/sec). Due to the complexity of rivers and the channel network of rivers, it is difficult to identify the sources of nutrients, especially the sources of nitrogen and its biogeochemical processes in the Taihu Lake watershed. To our knowledge, few works on N isotopes have been completed in the context of the region, probably resulting from the fact that isotopic signatures of N sources in such regions are obscured by complicated hydrographic connections and the mixing of multiple N sources with overlapping isotopic composition as well as *in situ* biogeochemical processes.

The present study analyzed the chemical parameters and isotopic compositions of NO_3^- , NH_4^+ , and particulate organic

matter (POM) in the water in the plain river network region. The specific goals of the study were to (1) evaluate the spatial variations in dissolved inorganic nitrogen (DIN) concentrations and isotopic compositions as well as the dominant sources of river NO_3^- , (2) evaluate the relative role of NO_3^- and NH_4^+ in the origin of POM in the river using DIN isotopic data, and (3) investigate possible seasonal shifts in nutrient sources (NO_3^- or NH_4^+) for phytoplankton in the river.

1. Materials and methods

1.1. Study site

Known as the main channel between Gehu Lake and Zhushan Bay, the Taige River empties into the northern part of Zhushan Bay of Taihu Lake (Fig. 1). The river is about 22.4 km long and 40 m wide on average. The slope ratio is 1:3. The mean annual discharge and velocity are $10.8 \text{ m}^3/\text{sec}$ and 0.13 m/sec, respectively. Fourteen main tributaries are directly connected to the Taige River. At most times, the flow direction is from north to south, but the direction is occasionally unsteady because of the variation in water levels and sluice gates of the tributaries. The river catchment is located in the Wujin District in Jiangsu Province. This area, with a population of about 0.18 million people and an area of 179.06 km^2 , is a typical plain river network region in the east of China, where the economy is developed and water pollution is severe. The area is a mixed land use watershed, which mainly serves agriculture (38% of land area in crops), residential, and industrial purposes. Many villages and townships characterized by high population density, urbanization, and economic development are situated along the banks of the upstream water catchment area of the Taige River. Water pollutants in the Taige River originate from different pollution sources and combinations of pollutants.

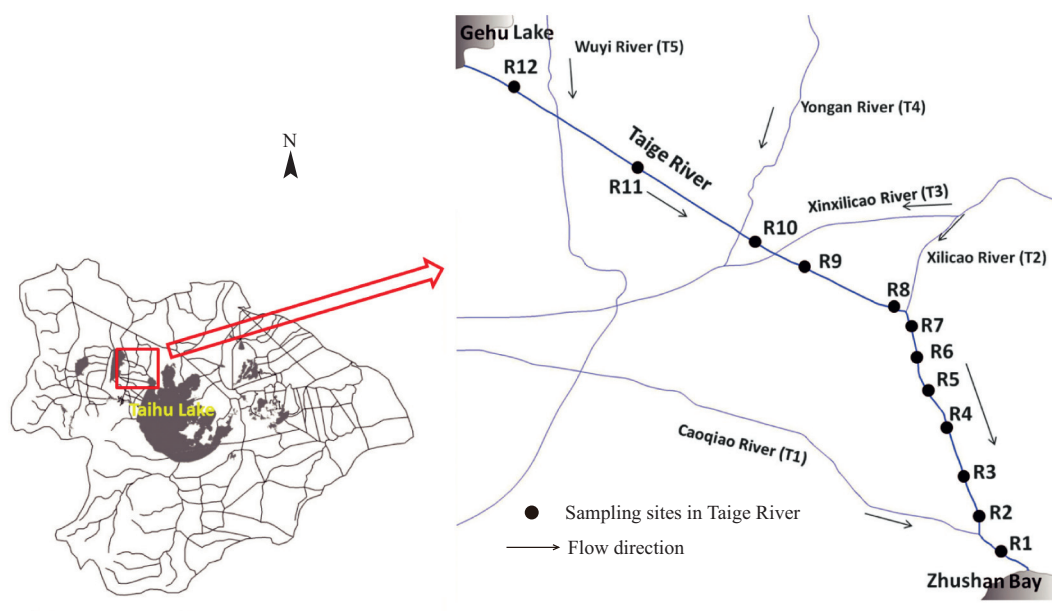


Fig. 1 – Sampling sites in the main stream of the Taige River.

1.2. Sampling

Twelve sites in the main stream of the Taige River were chosen for water sample collection (Fig. 1). Samplings in the main stream were performed on 9 January, 6 March, 7 May and 9 July in 2013. The discharge and flow velocity in each site was determined using an Acoustic Doppler Current Profiler. All sampling equipment was pre-cleaned with deionized water. Water temperature and dissolved oxygen (DO) were measured using a portable instrument (YSI 650MDS, Xylem, Yellow Springs, Ohio, USA) in situ. Water samples at 20 cm depth in the middle of the river were collected by boat and stored in pre-cleaned polyethylene bottles. About 300 mL of collected water samples was filtered on site through pre-combusted (450°C, 12 hr) and individually pre-weighed glass fiber filters (Whatman, GF/F, 47 mm in diameter). Filtrate and filter samples were stored at –30°C until needed for further treatments. In addition, 0.6 L of water samples was immediately carried back to the laboratory for nutrient concentration determination, within 24 hr at most.

1.3. Analytical methods

Water samples were filtered through 0.45 µm cellulose-acetate filter paper into polyethylene bottles and stored at 4°C until analysis. $\text{NH}_4^+\text{-N}$ concentrations were analyzed on a continuous flow analyzer. $\text{NO}_3^+\text{-N}$, nitrite ($\text{NO}_2^+\text{-N}$), and chloride analyses were carried out by ion chromatography.

Isotopic analyses of N and oxygen of NO_3^- were carried out using the denitrifier method (Sigman et al., 2001; Casciotti et al., 2002) based on the isotopic analysis of the nitrous oxide (N_2O) produced by denitrifying *Pseudomonas aureofaciens*. The N_2O was concentrated and purified on a Tracegas system, and the isotopic composition was determined using an isotope ratio mass spectrometer (IRMS; Isoprime100, Isoprime, Cheadle, UK) calibrated with ultra-high purity N_2 gas against air N (Xu et al., 2013). During measurement of $\delta^{15}\text{N}$, USGS34 potassium nitrate (KNO_3 , $\delta^{15}\text{N} = -1.8\text{‰} \pm 0.2\text{‰}$) and USGS32 KNO_3 ($\delta^{15}\text{N} = +18\text{‰} \pm 1\text{‰}$) standards were used to correct the values obtained. For $\delta^{18}\text{O}$ measurement, the samples were referenced using USGS34 ($\delta^{18}\text{O} = +27.9\text{‰} \pm 0.6\text{‰}$) and USGS35 sodium nitrate (NaNO_3 , $\delta^{18}\text{O} = +57.5\text{‰} \pm 0.6\text{‰}$) standards. The stable isotope ratios are expressed in delta (δ) units and a per mil (‰) notation relative to an international standard:

$$\delta_{\text{sample}} = (R_{\text{sample}}/R_{\text{standard}} - 1) \times 1000\text{‰} \quad (1)$$

where, R_{sample} and R_{standard} are the $^{15}\text{N}/^{14}\text{N}$ or $^{18}\text{O}/^{16}\text{O}$ ratios of the sample and standard for $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$, respectively. Values of $\delta^{15}\text{N}$ are reported relative to N_2 in atmospheric air, and $\delta^{18}\text{O}$ values are reported relative to Vienna Standard Mean Ocean Water. The analytical precision of the $\delta^{15}\text{N}\text{-NO}_3^-$ and $\delta^{18}\text{O}\text{-NO}_3^-$ are $\pm 0.2\text{‰}$ and $\pm 0.5\text{‰}$, respectively.

Determination of $\delta^{15}\text{N}\text{-NH}_4^+$ values was carried out using the ammonia distillation method of Velinski et al. (1989). The analytical precision of the $\delta^{15}\text{N}\text{-NH}_4^+$ samples was $\pm 0.5\text{‰}$. For the determination of the POM concentration as well as the $\delta^{13}\text{C}\text{-POC}$ and $\delta^{15}\text{N}\text{-PN}$ values, the GF/F filters were treated as follows in the laboratory. The filter samples were slightly wetted by hydrochloric acid (0.6 mol/L) and put in a desiccator

overnight with hydrochloric acid (0.6 mol/L) fumes to remove inorganic carbon. Subsequently, the filters were rinsed with deionized water to remove chloride, dried using a vacuum freeze dryer for 24 hr, and weighed to calculate POM concentration by weight difference. Then, $\delta^{13}\text{C}\text{-POC}$, $\delta^{15}\text{N}\text{-PN}$, and C/N ratios were measured in an elemental analyzer (vario PYRO cube, Elementar, Hanau, Germany) connected online to an IRMS. Values are reported relative to atmospheric N_2 ($\delta^{15}\text{N}$) and Vienna PeeDee Belemnite ($\delta^{13}\text{C}$). The analytical precision for both stable isotope ratios was $\pm 0.2\text{‰}$.

2. Results and discussion

2.1. Hydrological features and seasonal variation of water chemistry

The flow quantity (Q) and velocity (V) of the Taige River in the study period were $21.75 \pm 13.42 \text{ m}^3/\text{sec}$ and $0.28 \pm 0.16 \text{ m/sec}$ (mean \pm standard deviation (SD)), respectively. The factor that significantly influenced the hydrological condition of the river was the flow input from the Xilicao River (T2; $Q = 24.30 \pm 8.75 \text{ m}^3/\text{sec}$), which divides the Taige River into two sections with different hydrological features (Fig. 2). The Q of the upstream (before site R8), with lower Q ($7.17 \pm 2.52 \text{ m}^3/\text{sec}$) and V ($0.11 \pm 0.05 \text{ m/sec}$) than the downstream ($Q = 33.90 \pm 0.57 \text{ m}^3/\text{sec}$, $V = 0.42 \pm 0.04 \text{ m/sec}$), was strongly influenced by the Wuyi River (T5), Yongan River (T4), and Xinxilicao River (T3). In the downstream, the Q was almost constant and the V decreased slowly because of the backwater region of Zhushan Bay.

During the sampling period, $\text{NH}_4^+\text{-N}$, $\text{NO}_3^+\text{-N}$, $\text{NO}_2^+\text{-N}$ and Cl^- concentrations ranged from 0.35 to 4.01 mg/L, from 0.42 to 3.75 mg/L, from 0.06 to 0.32 mg/L and from 56.1 to 130.2 mg/L, respectively. In order to get a better understanding of the temporal variability of water chemistry in the river water, it was helpful to divide the data into two seasonal periods, winter (January and March, average temperature = 9.2°C) and summer (May and July, average temperature = 25.5°C). $\text{NH}_4^+\text{-N}$, $\text{NO}_3^+\text{-N}$ and Cl^- concentrations in winter were higher than in summer, while NO_2^- concentrations were higher in summer

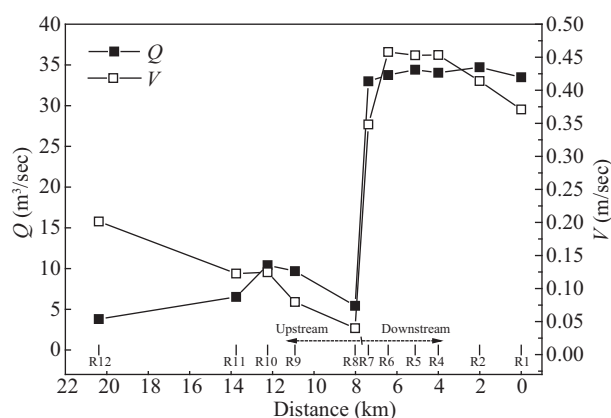


Fig. 2 – Longitudinal plots of averaged flow quantity (Q) and velocity (V) in the Taige River referenced to kilometers downstream from site R1.

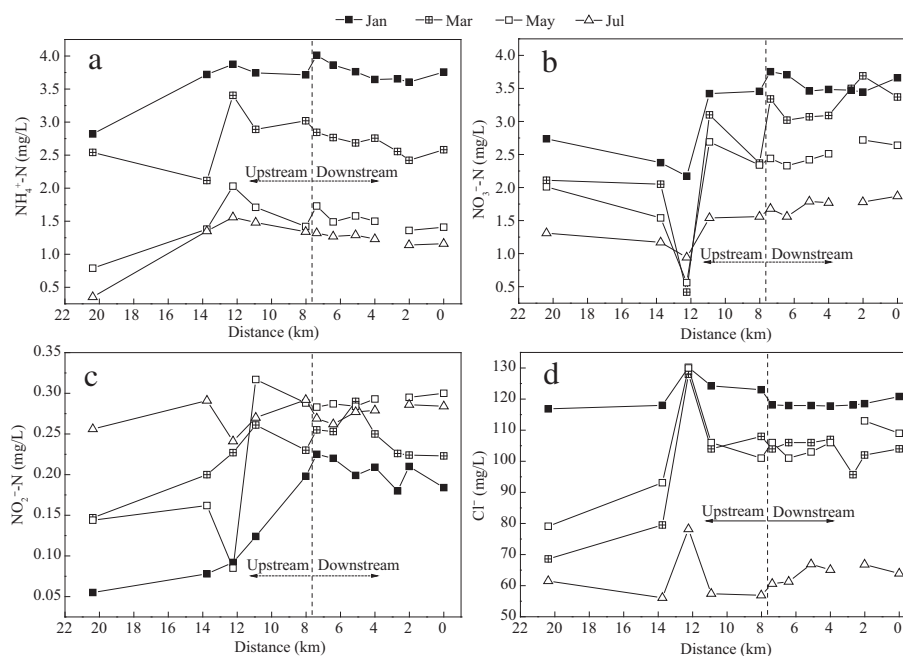


Fig. 3 – Longitudinal plots of (a) $\text{NH}_4^+\text{-N}$, (b) $\text{NO}_3^-\text{-N}$, (c) $\text{NO}_2^-\text{-N}$, and (d) Cl^- concentrations referenced to kilometers downstream from site 1. $\text{NH}_4^+\text{-N}$: ammonium; $\text{NO}_3^-\text{-N}$: nitrate nitrogen; $\text{NO}_2^-\text{-N}$: nitrite nitrogen; Cl^- : chloride.

(Fig. 3). This finding can be attributed to the dilution by the rain in the rainy season (June to July) in the study area.

Following the trends of the hydrological features, $\text{NH}_4^+\text{-N}$, $\text{NO}_3^-\text{-N}$, $\text{NO}_2^-\text{-N}$ and Cl^- concentrations in the upstream varied more extensively than in the downstream. At site R10 in the upstream, $\text{NH}_4^+\text{-N}$ and Cl^- concentrations increased and $\text{NO}_3^-\text{-N}$ concentration decreased sharply because of the input water of the Yongan River (T4), which drained urbanized areas (the city of Changzhou) and was polluted by domestic sewage and livestock wastewater with additional high $\text{NH}_4^+\text{-N}$ and Cl^- concentrations (Fig. 3a, b and 3d). The mean concentration of DO measured at Yongan River was 1.2 mg/L, which was ideal for denitrification (DO concentration <1 mg/L to 2 mg/L) (Rivett et al., 2008) and caused the low content of $\text{NO}_3^-\text{-N}$ in the Yongan River and in R10. In the downstream, $\text{NH}_4^+\text{-N}$ concentrations decreased slowly, while NO_3^- concentrations increased. The Cl^- concentrations had little change.

2.2. Spatial and temporal isotopic composition of NO_3^- , NH_4^+ and POM

The isotopic compositions of the NO_3^- in the Taige River were found to range from +4.56‰ to +8.79‰ ($n = 45$, mean value = +6.83‰) for $\delta^{15}\text{N}$ and from −0.24‰ to +8.17‰ ($n = 44$, mean value = +3.18‰) for $\delta^{18}\text{O}$. The $\delta^{15}\text{N}\text{-NH}_4^+$ and $\delta^{15}\text{N}\text{-POM}$ values ranged from −0.96‰ to +12.82‰ ($n = 45$, mean value = +4.58‰) and from −3.83‰ and +0.66‰ ($n = 44$, mean value = −1.46‰), respectively. The $\delta^{15}\text{N}$ and $\delta^{18}\text{O}\text{-NO}_3^-$ values in winter (mean \pm SD = +7.36‰ \pm 0.98‰, +4.66‰ \pm 1.25‰, respectively) were higher than those in summer (mean \pm SD = +6.23‰ \pm 1.02‰, +1.39‰ \pm 1.02‰, respectively) (Fig. 4a and b). The $\delta^{15}\text{N}\text{-POM}$ values were also higher in winter (Fig. 4d), but the $\delta^{15}\text{N}\text{-NH}_4^+$ values in winter (mean \pm SD = +2.60‰ \pm 1.84‰) were lower than those in summer (mean \pm

SD = +6.85‰ \pm 2.59‰) (Fig. 4c). The reasons for temporal variation of isotopic composition are discussed in Section 2.3.

Longitudinal variability in isotopic composition in the Taige River was significantly controlled by shifting hydrologic conditions. The isotopic composition of NO_3^- , NH_4^+ and POM in the upstream varied more extensively because of the inputs of various pollution sources with different isotopic composition in tributaries in the upstream. The Yongan River (T4), which was seriously polluted with high NH_4^+ and low DO content (see Section 2.1), had high $\delta^{15}\text{N}$ and $\delta^{18}\text{O}\text{-NO}_3^-$ values (mean \pm SD = +16.21‰ \pm 3.81‰, +11.00‰ \pm 3.09‰, respectively). The input of T4 resulted in the increased $\delta^{15}\text{N}$ and $\delta^{18}\text{O}\text{-NO}_3^-$ values at the site R10 in the Taige River (Fig. 4a and b). At the site R9, the $\delta^{15}\text{N}\text{-NH}_4^+$ and $\delta^{15}\text{N}\text{-POM}$ values increased sharply and then decreased (Fig. 4c and d). This was caused by the Xinxilicao River (T3) input with high $\delta^{15}\text{N}\text{-NH}_4^+$ (mean \pm SD = +9.40‰ \pm 3.81‰) and $\delta^{15}\text{N}\text{-POM}$ values (mean \pm SD = −0.88‰ \pm 0.59‰). In the downstream of the Taige River, the $\delta^{15}\text{N}\text{-NO}_3^-$, $\delta^{18}\text{O}\text{-NO}_3^-$ and $\delta^{15}\text{N}\text{-POM}$ values showed a clear increasing tendency (The reasons are discussed in Section 2.3) (Fig. 4a, b and d), while $\delta^{15}\text{N}\text{-NH}_4^+$ values showed no such clear development (Fig. 4c).

2.3. Nitrate source identification determined by dual-isotopic approach

The following sources possibly contributed to the NO_3^- budget in the Taige River in winter: precipitation, chemical fertilizer, nitrification of N-containing organic materials, livestock, and urban sewage effluent. The boxes in Fig. 5 show typical ranges of $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ for various natural and anthropogenic sources of NO_3^- .

The contribution of NO_3^- from precipitation was probably negligible for the waters of the Taige River considering the low $\delta^{15}\text{N}\text{-NO}_3^-$ (−13‰ to +13‰, close to 0‰) and high $\delta^{18}\text{O}\text{-NO}_3^-$

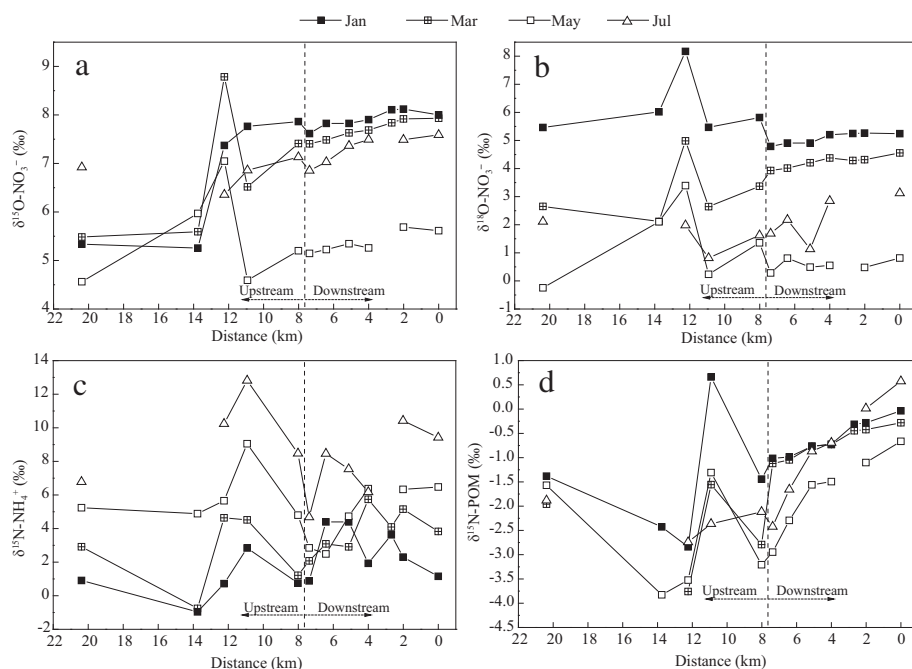


Fig. 4 – Longitudinal plots of (a) $\delta^{15}\text{N-NO}_3^-$, (b) $\delta^{18}\text{O-NO}_3^-$, (c) $\delta^{15}\text{N-NH}_4^+$, and (d) $\delta^{15}\text{N-POM}$ values referenced to kilometers downstream from site 1. $\text{NH}_4^+\text{-N}$: ammonium; $\text{NO}_3\text{-N}$: nitrate nitrogen; POM: particulate organic matter.

values (+25‰ to +75‰) in rain (Wassenaar, 1995; Durka et al., 1994; Kendall, 1998). In this district, fertilizer N was almost all in its reduced form, i.e., urea and ammonium, and the NO_3^- form was negligible. Ammonium fertilizer was not directly distributed into the rivers at the time of sampling because of its low $\delta^{15}\text{N-NO}_3^-$ values (−6‰ to +6‰). However, the isotopic composition of NO_3^- from chemical fertilizer would be

modified by biological processes during the flow of water into rivers from the soil and unsaturated zone (Mengis et al., 2001). Manure and sewage are enriched with ^{15}N relative to other N sources, and the $\delta^{15}\text{N}$ values of NO_3^- range from +4‰ to +25‰ (Xue et al., 2009). In this study, the $\delta^{15}\text{N-NO}_3^-$ and $\delta^{18}\text{O-NO}_3^-$ values of the water samples were in the range of “soil N” and “manure and sewage” source boxes, thus

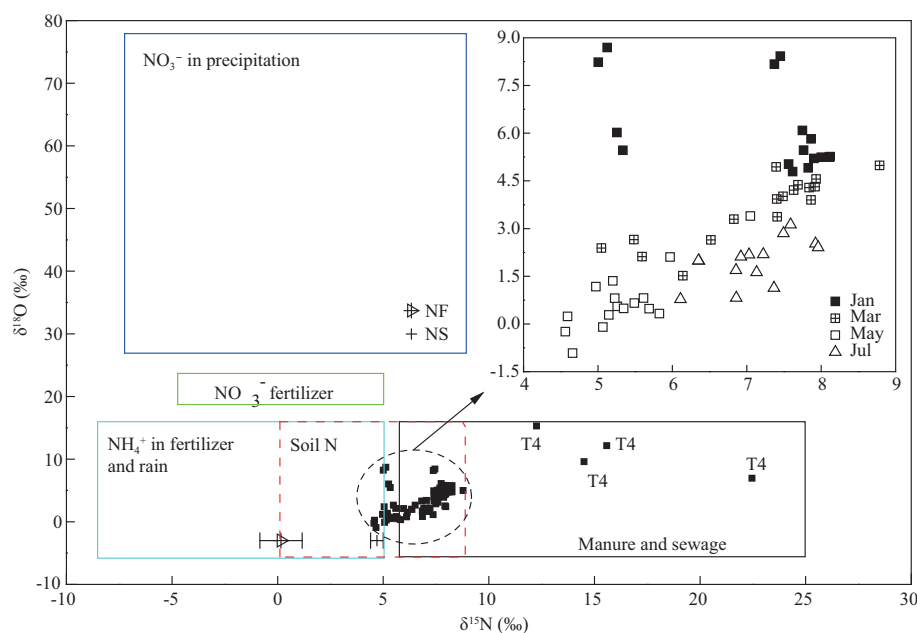


Fig. 5 – Cross plots of $\delta^{18}\text{O-NO}_3^-$ versus $\delta^{15}\text{N-NO}_3^-$ in water samples. The box diagram shows the isotopic composition of various sources (adapted from Kendall, 1998; Xue et al., 2009; Nestler et al., 2011). NS represents NO_3^- derived from soil organic matter nitrification; NF represents NO_3^- derived from chemical fertilizer nitrification. The $\delta^{15}\text{N}$ and $\delta^{18}\text{O-NO}_3^-$ values of NS and NF are expressed by mean values with error bars (Ding et al., 2014). NO_3^- : nitrate.

indicating that manure, urban sewage effluent, and nitrification of N-containing organic matter (including “modified fertilizer”) were likely the major sources of NO_3^- in the Taige River. These results are in good agreement with those by Townsend-Small et al. (2007), Chen et al. (2012), and Xing et al. (2001), who used the NO_3^- isotopic method and found that sewage-derived N makes a dominant contribution to the rivers to the northwest of Taihu Lake.

The $\delta^{15}\text{N}$ and $\delta^{18}\text{O}-\text{NO}_3^-$ values in the Yongan River (T4) were located in the “manure and sewage” source box. This was due to wastewater discharge from many livestock and poultry farms along the riverside of the Yongan River, and these NO_3^- sources are known to have high $\delta^{15}\text{N}-\text{NO}_3^-$ values that can reach +14‰ and sometimes even +20‰ (Heaton, 1986). Furthermore, other biochemical processes (mostly denitrification) in the Yongan River could also elevate the $\delta^{15}\text{N}$ and $\delta^{18}\text{O}-\text{NO}_3^-$ values.

Fig. 5 also indicates that the $\delta^{15}\text{N}$ and $\delta^{18}\text{O}-\text{NO}_3^-$ values in winter were higher than those in summer. In winter, the low flow and anthropogenic inputs (manure and urban sewage effluent) resulted in more serious NO_3^- pollution in the Taige River (Fig. 4b). This was supported by NO_3^- -N versus Cl^- concentration relationships in the Taige River. Chloride is a good indicator of sewage impact and dilution because it is not subject to physical, chemical, and biological processes (Liu et al., 2006). High Cl^- content mainly resulted from domestic sewage and livestock wastewater. NO_3^- -N versus Cl^- concentrations in river water showed that river water was more significantly affected by manure and urban sewage effluent, with higher $\delta^{15}\text{N}$ and $\delta^{18}\text{O}-\text{NO}_3^-$ values in winter than in summer (Fig. 6). In summer, the $\delta^{15}\text{N}$ and $\delta^{18}\text{O}-\text{NO}_3^-$ values of the Taige River were close to the low $\delta^{15}\text{N}$ and $\delta^{18}\text{O}-\text{NO}_3^-$ values of soil organic matter and chemical fertilizer in the study area (Fig. 5). This indicated that the intensified microbial nitrification of the soil organic matter and chemical fertilizer resulted in low $\delta^{15}\text{N}$ and $\delta^{18}\text{O}-\text{NO}_3^-$ in river water as climate warming and fertilizing activities took place in summer. Furthermore, dilution caused by precipitation in the rainy season could also result in low isotopic composition for NO_3^- in river water.

2.4. Assessment of the transformations of nitrogen in the downstream

According to the distributions of hydrological and water chemistry features, the mixing process of the three guest

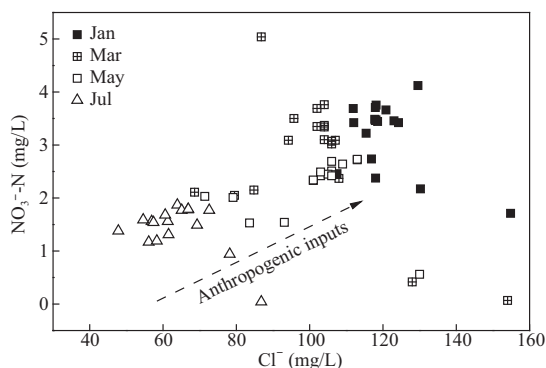


Fig. 6 – NO_3^- -N versus Cl^- concentrations in river water. NO_3^- -N: nitrate nitrogen; Cl^- : chloride.

rivers (Wuyi River, Yongan River, and Xinxiliao River) with different $\delta^{15}\text{N}$ and $\delta^{18}\text{O}-\text{NO}_3^-$ values caused the distinction between isotopic compositions in the upstream and the downstream of the Taige River. To further understand the biogeochemical N process in the Taige River, the following analysis focuses on the downstream of the Taige River, which has only minor variations in hydrological conditions.

During the sampling period, a decrease in NH_4^+ -N concentrations was observed in the downstream (Fig. 3a). This decrease could be attributed to ammonia volatilization, nitrification, and uptake by biota. The $\delta^{15}\text{N}-\text{NH}_4^+$ values in the downstream were small and could not be interpreted as the result of a ^{15}N enrichment of residual NH_4^+ resulting from volatilization (Chang et al., 2002; Sebilo et al., 2006). Either microbial nitrification or NH_4^+ assimilation by algae could result in increased $\delta^{15}\text{N}-\text{NH}_4^+$ values and decreased NH_4^+ -N concentration. The negative correlation between $\delta^{15}\text{N}-\text{NH}_4^+$ values and the natural logarithm of NH_4^+ -N concentration in the downstream indicated that nitrification and assimilation probably occurred simultaneously in the downstream of the Taige River (Fig. 7).

Although NH_4^+ assimilation in aquatic algae occurred (as discussed below), we believe that nitrification played a more important role in the decreasing NH_4^+ -N concentrations. First, coupled with the decreasing NH_4^+ -N concentrations, the NO_3^- -N concentrations increased in the downstream, indicating the occurrence of nitrification (Fig. 3a, b). Furthermore, in our study, the N isotopic enrichment factors (ϵ) during ammonium utilization in the downstream ranged from –13.88‰ in March to –29.00‰ in July (except in January, when the water temperature was low and microbial nitrification was restricted, $r^2 = 0.09$, $p > 0.05$) (Fig. 7). These factors (ϵ) were in the range of the N isotopic enrichment factors (ϵ) during nitrification (–12‰ to –29‰) (Kendall, 1998).

To better understand the nitrogen assimilation in the Taige River, we analyzed not only the relationship between $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ of NO_3^- but also the relationships between $\delta^{15}\text{N}-\text{POM}$ and isotopic compositions of DIN.

In addition to the increase in NO_3^- $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values (Fig. 4a, b), we found a significant correlation between NO_3^- $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values in the downstream in January ($r^2 = 0.81$, $p < 0.01$) and in March ($r^2 = 0.75$, $p < 0.01$) (Fig. 8). Algae/biota

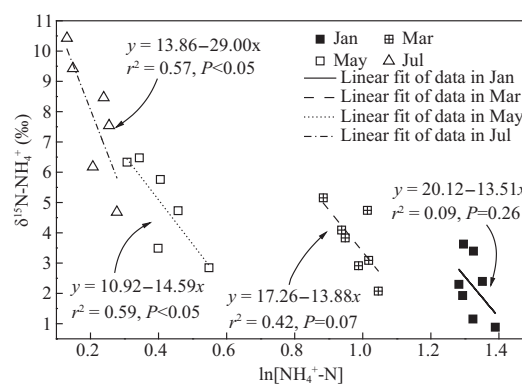


Fig. 7 – Dependence of $\delta^{15}\text{N}-\text{NH}_4^+$ values on the natural logarithm of NH_4^+ -N concentration ($\ln[\text{NH}_4^+-\text{N}]$) in the downstream of the Taige River, with the lines indicating linear regression fitting. NH_4^+ -N: ammonium.

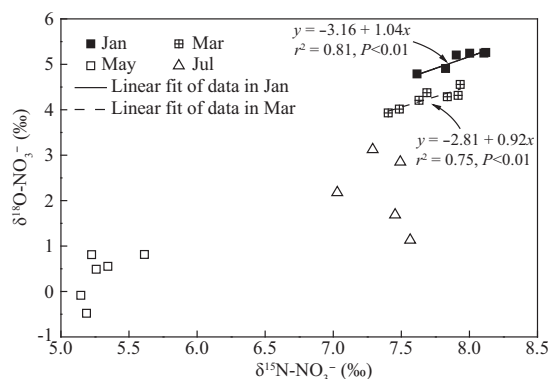


Fig. 8 – $\delta^{15}\text{N}\text{-NO}_3^-$ versus $\delta^{18}\text{O}\text{-NO}_3^-$ in the downstream water samples from the Taige River shown by a linear regression line. NO_3^- : nitrate.

generally prefer the uptake of light isotopes of NO_3^- , which would lead to enrichment with heavy isotopes in residual NO_3^- (Battaglin et al., 2001; Johannsen et al., 2008). The ratio of the increase in $\delta^{18}\text{O}$ and $\delta^{15}\text{N}$ of river NO_3^- in the downstream was 1.04 in January and 0.92 in March, which were closer to the ~1:1 increase reported for NO_3^- assimilation of marine phytoplankton (Granger et al., 2004) than to the 1:1.3 or the 1:2.1 increase reported for denitrification in freshwaters (Böttcher et al., 1990; Kendall, 1998; Mengis et al., 1999; Burns et al., 2009; Kaushal et al., 2011). The results indicated that NO_3^- assimilation by phytoplankton was responsible for the increasing $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values of NO_3^- in winter.

In the study area, many laboratory and field studies have indicated that the dominant N source responsible for the growth of phytoplankton was NH_4^+ (McCarthy et al., 2007; Paerl et al., 2011; Zhou et al., 2013). Our study demonstrates that the presence of POM in the Taige River was mainly derived from aquatic phytoplankton without terrigenous organic matter playing a role (Fig. 9). Thus, the relationships between $\delta^{15}\text{N}\text{-POM}$ and isotopic compositions of DIN could help us to understand the nitrogen assimilation. Fig. 10a shows two positive linear correlations existing between $\delta^{15}\text{N}\text{-POM}$ and $\delta^{15}\text{N}\text{-NO}_3^-$ values in winter, which supported NO_3^- assimilation

by phytoplankton in winter, as discussed before using NO_3^- $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values. Three other positive linear correlations were found between $\delta^{15}\text{N}\text{-POM}$ and $\delta^{15}\text{N}\text{-NH}_4^+$ values from March to July (Fig. 10b), which indicated that NH_4^+ assimilation in aquatic algae occurred and was partly responsible for the decreasing $\text{NH}_4^+\text{-N}$ concentrations. Therefore, we concluded that the phytoplankton in the Taige River probably utilized NH_4^+ preferentially and mainly in summer, while in winter, NO_3^- assimilation by phytoplankton was dominant. The results were supported by a laboratory study and field monitoring in Taihu Lake conducted by Zhou et al. (2013). They found *Microcystis* always assimilated NH_4^+ preferentially, but when $\text{NH}_4^+\text{-N}$ concentration exceeded 2 mg/L, their growth rate declined sharply. In our study, $\text{NH}_4^+\text{-N}$ concentrations of the Taige River were above 2 mg/L in winter and were below 2 mg/L in summer (Fig. 3a), which could partly explain the assimilation of different nitrogen forms in winter and in summer. Further systematic study should be performed to clearly understand N cycling in the area.

3. Conclusions

In our study of the Taige River, we found the river was polluted more seriously by anthropogenic inputs in winter than in summer, based on DIN concentrations. Manure and urban sewage effluent, as well as the nitrification of N-containing organic materials, were the major sources of N. Efforts to reduce N exported from livestock waste, urban wastewater and soil should be a priority. In the downstream of the Taige River, nitrification and assimilation probably occurred simultaneously. Nitrification played a more important role than assimilation in the decreasing $\text{NH}_4^+\text{-N}$ concentrations. Nitrification strengthened as the water temperature increased. Based on the data of the isotopic composition of DIN and POM, we also found that the phytoplankton in the Taige River probably utilized NH_4^+ preferentially and mainly in summer, while in winter, NO_3^- assimilation by phytoplankton was dominant. Our study demonstrates that the dual-isotope approach ($\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ of NO_3^-), combined with analyses of NH_4^+ and POM isotopic compositions, is a helpful tool in investigating sources of NO_3^- and N transformation processes along rivers in the plain river network region.

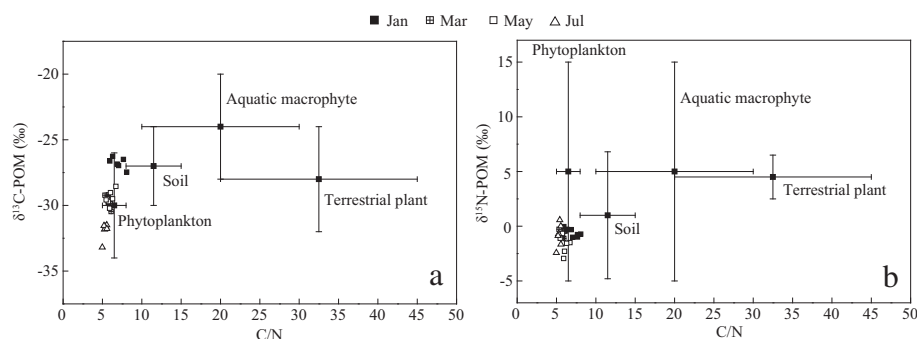


Fig. 9 – Relationships between (a) $\delta^{13}\text{C}\text{-POM}$ and C/N and (b) $\delta^{15}\text{N}\text{-POM}$ and C/N ratios in the Taige River, including the ranges of the potential sources of organic material.

Adapted from Hamilton and Lewis (1992), Angradi (1994), Thorp et al. (1998), Kendall (1998), Vuorio et al. (2006), Ogrinc et al. (2008), and Tang and Zhang (2010).

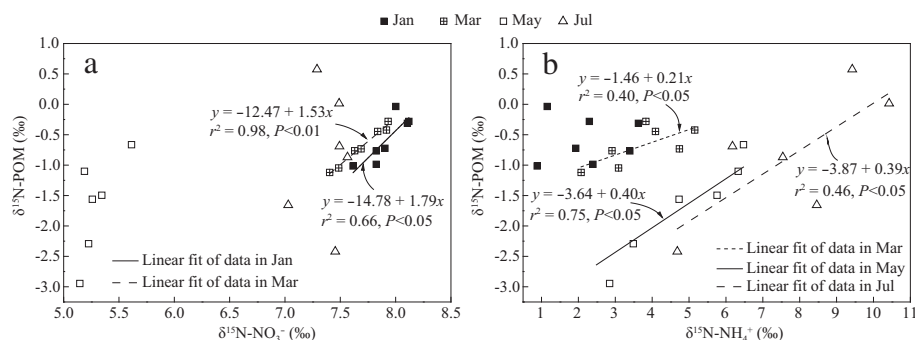


Fig. 10 – Relationships between (a) $\delta^{15}\text{N-POM}$ and $\delta^{15}\text{N-NO}_3^-$ values and (b) $\delta^{15}\text{N-POM}$ and $\delta^{15}\text{N-NH}_4^+$ values in the Taige River. POM: particulate organic matter; NO_3^- : nitrate; NH_4^+ : ammonium.

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